

## **Recent surges on Blomstrandbreen, Comfortlessbreen and Nathorstbreen, Svalbard**

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### **Introduction**

Very few glaciers in Svalbard are known to have shown more than one recorded surge (Hagen and others, 1993). A decreasing frequency of glacier surges in Svalbard was suggested by Dowdeswell and others (1995). Recent observations indicated a different situation and several new surges have been observed based on new interpretation and evidence (Sund and others, 2009). These results also showed that geometric changes were present several years prior to a visible indication of a surge following the pattern of previous descriptions.

Surge progress of three glaciers is described here (Fig 1): (1) Blomstrandbreen is a ~80 km<sup>2</sup> tidewater glacier draining into Kongsfjorden. The glacier is known to have surged around 1960 (Hagen and others, 1993). (2) Comfortlessbreen is ~60 km<sup>2</sup> and partly a tidewater glacier draining to Engelsbukta south of Ny-Ålesund. Sund and others (2009) reported a surge and advance of the glacier seen from (2006). Until then it had been unclear whether this was a surge-type glacier or not. (3) The tidewater glacier system in the inner part of Van Keulenfjorden is, due to former retreat, currently divided in two main branches: Nathorstbreen glacier system (NGS) and Liestølbreen – Doktorbreen. Liestøl (1977) suggested a maximum glacier extension by ~1870 probably caused by a surge, which was ~20 km further out the fjord compared to the situation in late 2008. Also Croot's (1988) interpretation of moraine patterns suggests several glaciers in the inner part of Van Keulenfjorden having gone through surges, however no years are specified. NGS consists of several glaciers where Dobrowolskibreen, the icefield Ljosfonn, Polakkbreen and Zawadzkiibreen are the most important ones, converging into Nathorstbreen draining out into the head of the fjord. Altogether they constituted some 390 km<sup>2</sup> in 2008. Results from Sund and others (2009) showed the development of NGS

through different stages resulting in a progressive advance of the terminus during winter 2008-09. These precursory observations will be completed in a future paper including a more detailed analysis of the surges.

## **Methods and Results**

We compared satellite images from Moderate-Resolution Imaging Spectroradiometer (MODIS - 250 m resolution), Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER-15 m resolution), Satellite Pour l'Observation de la Terre (SPOT-5 HRS - 5 m resolution), QuickBird panchromatic (0.6 m resolution) and recent aerial photos from the Norwegian Polar Institute (NPI), (0.6 m resolution). Front changes were mapped and some velocities derived from tracing of crevasses.

### *Blomstrandbreen*

During quiescence since the end of the last surge in 1966 (Hagen and others, 1993), the glacier has retreated 2.5 km and withdrawn from the island Blomstrandhalvøya, which previously was thought to be a peninsula. During the last years it has almost emerged above tidewater on the other side. Comparison of images acquired during summers 2007 (SPOT-5) and 2008 (QuickBird) shows no change of the glacier front, while there was a small advance between 2008 and 2009 (aerial photos). Although the change is not significant alone, also more important characteristics like new and increased crevassing up to 700 m a.s.l. at least by 2007 when compared to the situation on aerial images from 1990 (NPI) were found (Fig. 1). In 2009 some crevasses could be seen also in the uppermost part.

### *Comfortlessbreen*

At the centre of the tidewater terminus, Comfortlessbreen underwent a small retreat of ~250 m between 1990 and 2002, derived from aerial photos and ASTER. Subsequently, the changes of the front position from ASTER indicated a switch from quiescent to a more active phase already before 2004. Between 2002 and 2004 part of the front advanced up to 100 m. By 2009 (aerial images), the front had advanced 500-700 m since 2002 (Fig. 2), and now also the land based lobe was activated. For velocity measurements

we used aerial photos from end of July 2008 and a QuickBird image acquired one month later to trace crevasses. The QuickBird image was already orthorectified. From 14 aerial images (NPI), a new DEM was compiled and used for orthorectification. The velocities derived along almost the entire glacier, showed a one month average of  $\sim 2 \text{ m d}^{-1}$ , indicating block sliding, which is characteristic for surges. This occurred as the surface crevassing clearly increased (Fig. 3). As a comparison, the highest velocity from GNSS measurements carried out in April 2001 at a distance one third in length up from the terminus (Fig. 2) was  $0.45 \text{ m d}^{-1}$  (Z. Perski personal communication September, 2009). On the larger Kongsvegen in the same area maximum velocities are only  $\sim 5 \text{ m a}^{-1}$  during summer (Melvold and Hagen, 1998) while the maximum on e.g. Finsterwalderbreen is  $25 \text{ m a}^{-1}$  (Nuttall and others, 1997).

#### *Nathorstbreen glacier system*

Dobrowolskibreen was the first glacier in the system to surge (Sund and others, 2009), resulting in activation of the stagnant marginal ice at the northern side of the NGS front causing a small advance of this part. The SPOT-5 image from September 2008 proved there were no major changes at the terminus at this time. The last MODIS image from October 2008 confirms this. Although being at a coarser resolution, no changes are visible here. Thus the advance started during winter 2008-09, and by end of March 09, the terminus had advanced  $\sim 4.6 \text{ km}$ . By October 2009 the advance was minimum  $\sim 7.6 \text{ km}$  compared to the September 2008 position (Fig. 4).

The change of front positions induces an average velocity of  $\sim 20 \text{ m d}^{-1}$ , but because of calving the actual average velocity have been higher than this. The MODIS positions were also compared with front positions on photos taken from commercial airlines and correspond well to one another. In late August, major calving activity and rumble could be heard for several minutes at a distance of 15-20 km during fieldwork. Short movies based on time-lapse photographs showing parts of the surges of Comfortlessbreen and Nathorstbreen can be viewed on

[http://www.unis.no/35\\_staff/staff\\_webpages/geology/monica\\_sund/web/new/monica\\_sund\\_homepage.htm](http://www.unis.no/35_staff/staff_webpages/geology/monica_sund/web/new/monica_sund_homepage.htm)

## **Discussion**

It is approximately 50 years since the last surge of Blomstrandbreen. One requirement for a glacier to surge is sufficient built up (Meier and Post, 1969) which Blomstrandbreen, although a rather high estimated ELA of 500 m a.s.l. (Hagen and others, 1993), apparently has experienced. Svalbard glaciers generally experience slower initiations than other areas (e.g. Dowdeswell and others, 1991). On Comfortlessbreen, slight changes in front position several years prior to large and clearly visible surface changes and advance were observed. Several previous measurements of surge velocities in Svalbard are found to be at a level of 2-3 m d<sup>-1</sup> (Dowdeswell and Benham, 2003; Murray and others, 2003; Błaszczuk and others, 2009) and is in accordance with those observed on Comfortlessbreen. For NGS large elevation changes in the uppermost parts of the glaciers, of up to 70 m, took place several years prior to an extensive crevassing (Sund and others, 2009). The front velocities derived here were relatively high compared with many other Svalbard surges, and are more comparable to those recorded during the extensive surges of Negribreen and Bråsvellbreen in 1935-38 (Liestøl, 1969). The NGS surge is the largest observed in Svalbard since these surges.

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**Figures:**

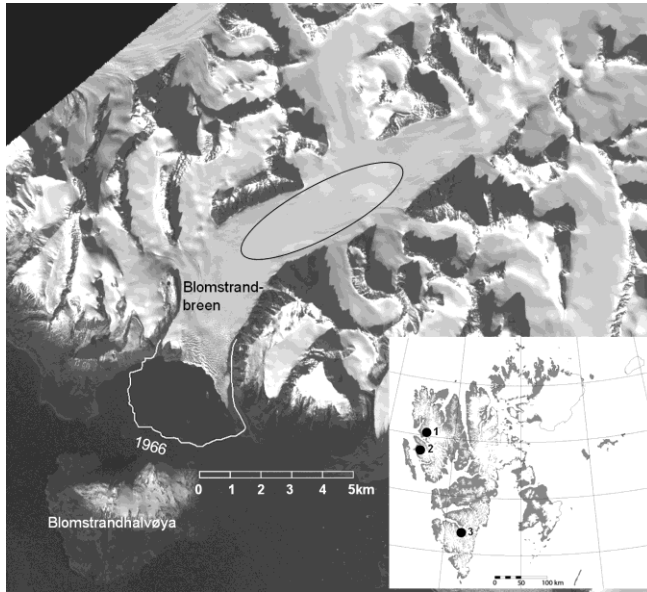


Fig. 1. Inset: map of Svalbard with location of Blomstrandbreen (1), Comfortlessbreen (2) and the Nathorstbreen Glacier System (3). The front of Blomstrandbreen has retreated 2.5 km across the sound since 1966. Area with pronounced new crevassing is indicated with an oval. Background image: SPIRIT Program © CNES 2008 (2008) and Spot Image (2007) all rights reserved.

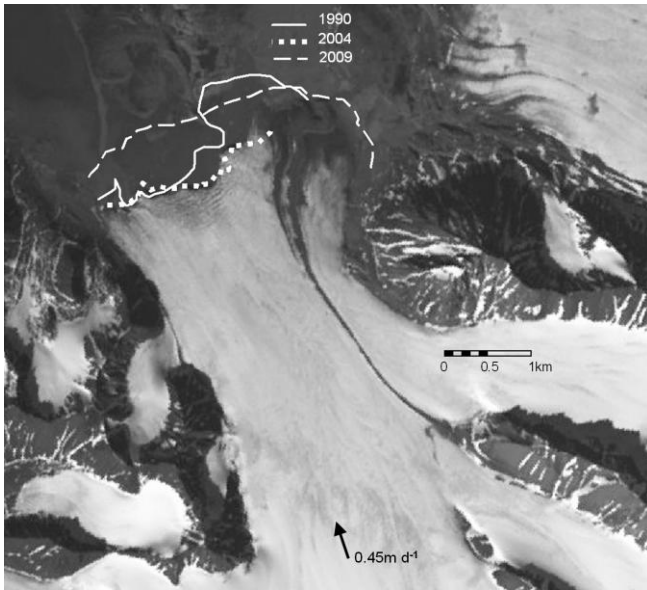


Fig. 2. Frontal changes of Comfortlessbreen based on aerial image from NPI (1990), ASTER (2004) and aerial photos from NPI (2009). GNSS velocity measured in 2001 (Z. Perski) indicated with an arrow. Background image: ASTER (2002).



Fig. 3. The upper part of Comfortlessbreen, showing characteristic transverse crevassing and sheared margins. The glacier flows towards photographer. Photo: M. Sund.



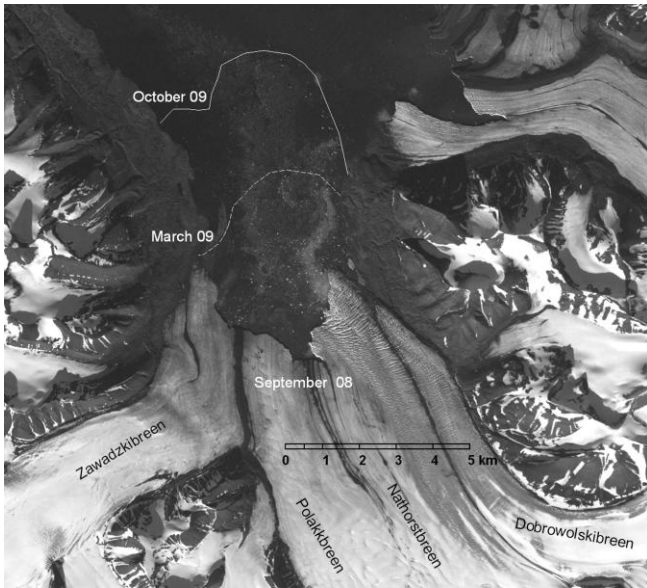


Fig. 4. Front positions and advance of the Nathorstbreen Glacier System during the period September 2008 to October 2009 based on MODIS images. Background image: SPOT-5. SPIRIT Program © CNES 2008 (2009) and Spot Image (2008) all rights reserved.